

The first high resolution diatom record from Lake Balaton, Hungary in Central Europe

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ABSTRACT

The first high resolution diatom record from Lake Balaton, Hungary in Central Europe

A high-resolution diatom record of the last 7500 years of Lake Balaton, a large, shallow calcareous lake in Central Europe is presented here. Diatom analyses were carried out on a 112 cm long core obtained from the Siófok Basin (eastern part of the lake) combined with loss-on-ignition, geochemical measurements and grain-size analysis. Altogether 140 diatom taxa were distinguished in this study; *Amphora pediculus*, *Cocconeis neothumensis*, *Fragilaria brevistriata* and *Karayevia clevei* were frequent and ubiquitous members of diatom assemblages. Fragilaroid taxa were dominant through the core in the shallow Lake Balaton, but episodes of higher ratios of planktic forms (*Lindavia balatonis*, *Aulacoseira granulata* and *Pantocsekiella ocellata*) before ca. 2000 cal yr BP, indicated high water level in the lake, and stronger fluctuation than exists today. The diatom-based reconstruction of lake level was in accordance with the geochemical and zoological records, proving the usability of siliceous remains in paleolimnological reconstructions in Lake Balaton, in spite of the unfavourable preservation conditions.

Key words: diatoms, geochemistry, Lake Balaton, lake level changes, shallow lake sediments

RESUMEN

El primer registro de diatomeas de alta resolución del lago Balaton, Hungría, en Europa Central

Aquí se presenta un registro de diatomeas de alta resolución del lago Balaton de los últimos 7500 años, un gran lago calcáreo poco profundo en Europa Central. Los análisis de diatomeas se llevaron a cabo en un testigo sedimentario de 112 cm de longitud obtenido en la cuenca Siófok (parte oriental del lago), combinado con pérdidas por ignición (LOI), mediciones geoquímicas y análisis de tamaño de grano. En este estudio se distinguieron 140 taxones de diatomeas; *Amphora pediculus*, *Cocconeis neothumensis*, *Fragilaria brevistriata* y *Karayevia clevei* fueron miembros frecuentes y ubicuos de los conjuntos de diatomeas. Los taxones de *Fragilaroides* eran dominantes a través del testigo en el somero lago Balaton, pero los episodios de mayor proporción de formas planctónicas (*Lindavia balatonis*, *Aulacoseira granulata* y *Pantocsekiella ocellata*) antes de ca. 2000 calyr BP, indicaron un nivel de agua alto en el lago y una fluctuación más fuerte que la que existe hoy en día. La reconstrucción del nivel del lago basada en diatomeas estuvo de acuerdo con los registros geoquímicos y zoológicos,

demostrando la utilidad de los restos silíceos en reconstrucciones paleolimnológicas en el lago Balaton, a pesar de las condiciones de preservación desfavorables.

Palabras clave: *diatomeas, geoquímica, lago Balaton, cambios en el nivel del lago, sedimentos someros del lago*

INTRODUCTION

Lakes are complex ecosystems, playing vital roles in regulating the global hydrological and biogeochemical cycles and acting as important parts of the global biosphere (Romero-Viana *et al.*, 2009; Williamson *et al.*, 2009). Paleolimnology – a rapidly developing field of science – gathers information about the ontogeny of different kinds of waterbodies, mires, wetlands, ponds and lakes. Among these, shallow lakes are extremely sensitive to any stressors, even as they carry exceptionally high biological conservation values (Kearns *et al.*, 1996, Heinsalu *et al.*, 2007b). Shallow lakes and wetlands are important subjects of hydrological reconstruction, as they are more vulnerable than deep lakes to water level fluctuation. Even a small scale variation can cause considerable changes due to their shallowness. It is especially true for such a large and shallow lake as Lake Balaton, where the wind, combined with shallowness makes the lake prone to re-sedimentation, affecting benthic life (Korponai *et al.*, 2010).

Diatoms are unicellular algae; they are sensitive indicators of environmental changes. They have been successfully used to monitor disturbances, like eutrophication, changes to the physical and chemical properties of water, as well as fluctuations of water level (Stoermer & Smol 1999). However, though there is a wide consensus about the usefulness of diatoms in quantitative paleolimnological reconstruction (Stoermer & Smol 1999; Romero-Viana *et al.*, 2009; Heinsalu *et al.*, 2007a, 2007b), in Lake Balaton, no high-resolution diatom record has been published so far; therefore, the study presented here is entirely new.

Paleolimnological research on Lake Balaton

In spite of the intensive and continuous long-term neolimnological studies of Lake Balaton (Hat-

vani *et al.*, 2011, Istvánovics & Honti, 2018), the paleolimnological exploration of the lake is sparse and uneven, and is sometimes even questioned (WEB1 2018). There has been no comprehensive modern survey conducted on the entire lake, and only limited efforts were made for using multi-proxy approaches for paleoecological reconstruction concerning the history of the lake. In particular, the possibly uneven and dissimilar development of the different basins of Lake Balaton has been neglected.

Nevertheless, from the end of the nineteenth century, the lake and its surroundings have been intensively studied (Lóczy, 1916). It must be mentioned about this period that the first diatom record by J. Pantocsek (1913) presented a very low-resolution list with presence/absence data of a 4 meter long core from Siófok Basin.

In 1981, the Geological Institute of Hungary (MÁFI) initiated a comprehensive survey of the limnogeology of Lake Balaton. Altogether 33 boreholes were drilled into the Quaternary layers down to 10–12 m. The thickness of the Holocene mud was 5 meters on average, but it was widely varying from place to place, e.g. 10 meters in the western basin and less than 2 meters in the eastern end of the lake (Cserny & Nagy-Bodor 2000, Cserny 2002, Tullner & Cserny, 2003). These studies revealed that Lake Balaton is 15 000–17 000 years old, and its development started with the formation of several shallow water lakes. Buczko *et al.* (2005) summarised the diatom data gathered by Márta Hajós working on 18 of these 33 cores. More than 350 taxa were listed in this compilation, but evaluation of data has never been completed. Shortcomings of the Márta Hajós dataset preclude it from being used in this study. No standard counting methods were followed (e.g. Battarbee, 1986), resulting in varying sample sizes, where the number of counted valves could range from less than ten to 1300 valves (Tó-29). Sometimes, even the counted valves of different slides were merged into one

single entry in the database. Consequently, there is no simple statistical method that can be applied for the Márta Hajós dataset. Additionally, in these 33 cores, the organic content of samples was not measured, leaving us without this standard geochemical proxy.

More recently Korponai *et al.* (2010) made the first attempt to use multi-proxy methods in the ancient Lake Balaton where extended wetlands were recently reconstructed (Hatvani *et al.*, 2011). The site for coring was chosen in a formerly water covered area on the south-western part of the lake. Three different stages were established concerning water level on the basis of geochemical data, subfossil Cladocera and diatom remains in the sediments of the Zalavár Pond, a part of the Kis-Balaton wetland (Buczko *et al.*, 2009; Korponai *et al.*, 2010). It should be emphasized that the coring site of Korponai *et al.* (2010) is beyond the present day open-water part of Lake Balaton, so the result of this multi-proxy study has its limitations concerning the history of the lake compared to the present coring site posi-

tioned centrally in a sedimentary basin which provides a sedimentary record embracing a longer time period.

Within the frame of the EuLakes project (WEB1 2018), there was an attempt to use high-resolution trophic reconstruction in the south-western basin of Lake Balaton. In conclusion, the applicability of diatom analyses was questioned for such a shallow, wind disturbed, calcareous lake as Lake Balaton (WEB1 2018).

The present study is, first of all, intended to test the applicability of diatoms as a paleolimnological proxy in the eastern basin of Lake Balaton. The goals of this work are twofold: first, to publish the first high-resolution diatom record from Lake Balaton with special attention to the preservation of small celled valves; secondly, to perform lake level reconstruction, using geochemistry, and subfossil diatom remains in the sediment. To achieve these goals, the Siófok Basin at the eastern end of Lake Balaton, characterized by high human pressure, but low direct input from the watershed, was chosen (Fig. 1.).

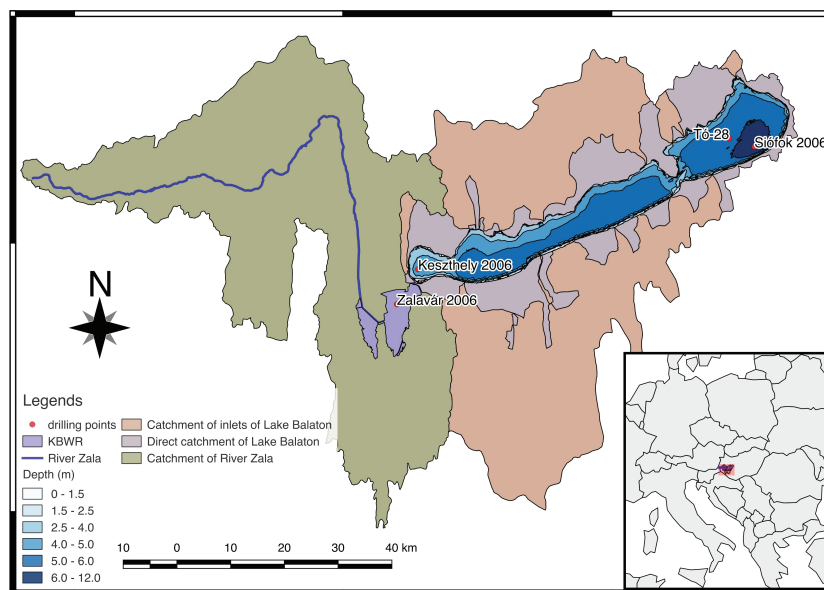


Figure 1. Lake Balaton and its catchment area. Note the four basins (Keszthely, Szigliget; Szemes, Siófok) and the difference between their catchment area size. Three cores were drilled in 2006 (Buczko *et al.*, 2009, Korponai *et al.*, 2010). Tó-28 drilled within the MÁFI actual geological survey (Cserny, 2002, Rokob *et al.*, 2008). *El Lago Balaton y su área de influencia. Tenga en cuenta las cuatro cuencas (Keszthely, Szigliget, Szemes, Siófok) y la diferencia entre el tamaño de su área de influencia. Se perforaron tres testigos sedimentarios en 2006 (Buczko et al., 2009, Korponai et al., 2010). Tó-28 perforado dentro de una verdadera investigación geológica MÁFI (Cserny. 2002, Rokob et al., 2008).*

MATERIAL AND METHODS

Study site

Lake Balaton, the largest shallow lake in Central Europe, is one of the best studied lakes in the world (Entz & Sebestyén 1946, Istvánovics *et al.*, 2007, Istvánovics & Honti 2018). Its catchment area is 5175 km², with a major inflow, the Zala River in the west, with an average discharge of about 5–10 m³/seg. The only outflow is the artificial Siófok canal constructed in the 1860s connecting the lake to the Danube River (Fig. 1). The Zala River enters Lake Balaton through the Kis-Balaton Wetland, a large area acting as a natural filter for suspended material and nutrients that would otherwise be deposited in the lake (Fig. 1. Hatvani *et al.*, 2011). The lacustrine sediments are rich in carbonate (calcareous mud) increasing from south-west (Keszthely) to north-east (Siófok) (Cserny & Nagy-Bodor, 2000; Tullner & Cserny, 2003).

Field and laboratory methods

A sediment core (SIOFOK2006) of 112 cm length was obtained from the deepest, central part of the Siófok Basin in Lake Balaton (Fig. 1.) with a 5 cm diameter Livingston piston sampler in May 2006. Subsamples were taken at every second cm for siliceous algae, grain size and geochemical analyses. The procedure for grain size determination started with disaggregation by shaking wet samples for 24 hours before fractionating them in a settling column. For loss-on-ignition analysis, samples were weighed and heated at 550 °C for 4 hours (Heiri *et al.*, 2001) and at 950 °C for 2 hours. Trace element content was determined by

ICP-OES - Thermo Iris Intrepid II XSP (Korponai *et al.*, 2010). For radiocarbon calibration of the core, macrophyte and charred plant remains were selected and (Table 1) analysed at the Laboratory of Environmental Studies, Institute of Nuclear Research, Hungarian Academy of Sciences in Debrecen. Dates were calibrated using the software CALIB REV5.0.0 (Stuiver *et al.*, 2005).

In order to analyse siliceous algae, sediment samples were prepared using standard digestion procedures (10 % hydrochloric acid, and 30 % H₂O₂ hot hydrogen-peroxide) (Battarbee, 1986), and at least 400 valves were counted in each sediment layer. For light microscope analysis, a LEICA DM LB2 was used (100 X HCX PLAN APO). Diatom identification followed the nomenclature of Hofmann *et al.* (2011) and Wetzel *et al.* (2015) for small celled naviculoids. Nomenclature was checked and updated according to AlgaeBase (Guiry & Guiry, 2016), accessed 2–25 April 2016, except genus *Pantocsekiella* where the concept of Ács *et al.* (2016) was followed. Stratigraphic zone boundaries were defined using CONISS on square-root transformed data without down-weighting of rare taxa with the Psimpoll 4.27 program (Bennett, 2005). The significant number of diatom assemblage zones (DAZ) was assessed by a broken-stick model implemented in “Rioja” package (Juggins, 2012).

RESULTS

Physical properties of the core

Lithology, age, organic material content expressed in loss on ignition (LOI) at 550 °C and at 950 °C, Ca and Mg content, Fe and Mn content and their ratio according to the grain size distribu-

Table 1. Results of the ¹⁴C dating measurements, SIOFOK2006 core; Lake Balaton. *Resultados de las dataciones con ¹⁴C, testigo SIOFOK2006; Lago Balaton.*

Lab-code	Depth below core top (cm)	D13C (PDB) ± 0.2 [‰]	¹⁴ C year BP	Cal. year (2s) Cal yr BP	Mid-point of 2sigma calibrated age range
deb-13938	Siófok 59-64 cm	-0.53	3920 ± 50	4468 – 4516	4492
deb-13947	Siófok 104-109 cm	+1.05	6330 ± 60	7158 – 7421	7394

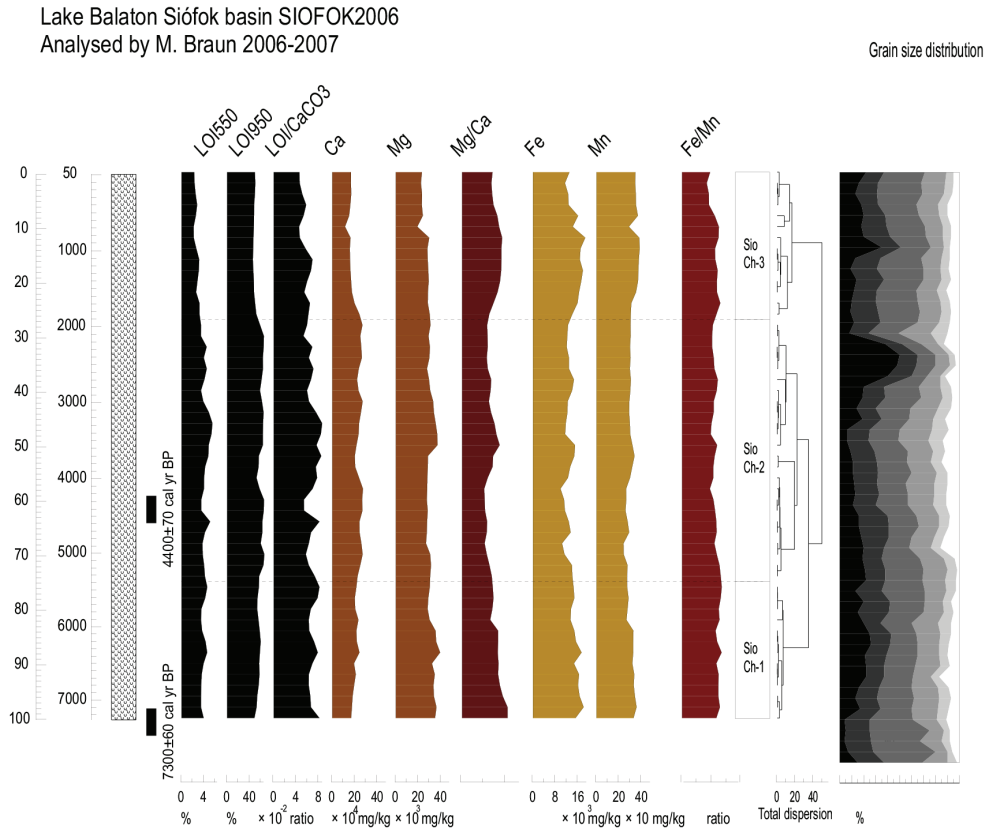


Figure 2. Physical properties of the core SIOFOK2006, Siófok Basin, Lake Balaton. Lithology, organic material content, main cations (Ca, Mg, Fe, Mn); Mg/Ca and Fe/Mn ratios are also plotted. Sio-Ch: local geochemical assemblage zones. Grain-size distribution, Sio-Gr: local grain size assemblage zones. *Propiedades físicas del testigo SIOFOK2006, Cuenca Siófok, Lago Balaton. Litología, contenido en materia orgánica, principales cationes (Ca, Mg, Fe, Mn); También se han representado las relaciones Mg/Ca y Fe/Mn. Sio-Ch: Zonas de agrupación geoquímica. Distribución granulométrica, Sio-Gr: Zonas granulométricas.*

tion are presented in Figure 2. Although the core SIOFOK2006 turned out to be quite uniform, homogeneous and unvaried in terms of lithology, and geochemistry, e.g. organic material content. Still, three zones could be distinguished.

Organic material content estimated by LOI₅₅₀ is low, varying in a very narrow range (2-5 %), and shows some decrease at the top of the core. The LOI₉₅₀ is high (40-65 %) and indicates three intervals. It is lowest at the top 30 cm. Higher values characterize the region between 30 and 75 cm, whereas below 75 cm has lower values. The topmost sample does not show high LOI values (LOI₅₅₀ 2 % and LOI₉₅₀ 46 %) indicating continuously disturbed, mixed sediments, preventing rich benthic vegetation on the bottom of the lake.

Five zones (Sio-Gr1–Sio-Gr5) were distinguished using the grain size distribution (Fig. 2). Generally, the core is silty, except for a sand fraction peaking between 28 and 38 cm, with grains of calcareous aggregates, most likely formed autogenetically as part of diagenetic processes.

The two successful measurements of radiocarbon dates are presented in Table 1. A linear sedimentation rate model was applied, establishing the linear relationship between age and depth. The high carbonate and low organic content complicated the radiocarbon dating of the core, making the applied time frame less than ideal concerning reliability. The studied 112 cm long core spans ca. 7500 years.

Table 2. List of diatom taxa in SIOFOK2006 core was found > 5 % in relative abundance in samples (%). The first column gives the taxon relative abundance in the whole core, while the second column show the number of occurrences in the 53 studied samples. *La lista de taxones de diatomeas en el testigo SIOFOK2006 fue > 5% en abundancia relativa en las muestras (%). La primera columna da la abundancia relativa del taxón en todo el testigo, mientras que en la segunda columna muestra el número de ocurrencias en las 53 muestras estudiadas.*

List of taxa in Siofók Basin of lake Balaton (V. Pozderka 2015-2016)	Relative abundance	Frequency
<i>Amphora copulata</i> (Kützing) Schoeman & R.E.M.Archibald	0.83	42
<i>Amphora pediculus</i> (Kützing) Grunow ex A.Schmidt	7.65	54
<i>Aneumastus minor</i> Lange-Bertalot	0.2	23
<i>Aneumastus tuscula</i> (Ehrenberg) D.G.Mann & A.J.Stickle in Round, Crawford & D.G.Mann	0.11	15
<i>Aulacoseira granulata</i> (Ehrenberg) Simonsen	6.95	54
<i>Belonastrium berlinensis</i> (Lemmermann) Round & Maidana	2.73	51
<i>Biremis zachariasii</i> (Reichelt) Edlund, N.A.Andresen & Soninkhisig	0.28	23
<i>Caloneis schumanniana</i> (Grunow) Cleve	0.15	5
<i>Campylodiscus hibernicus</i> Ehrenberg	0.12	14
<i>Cavinula jaernefeltii</i> (Hustedt) D.G.Mann & A.J.Stickle in Round. Crawford & D.G.Mann	0.1	6
<i>Cavinula scutelloides</i> (W.Smith) Lange-Bertalot in Lange-Bertalot & Metzeltin	0.46	39
<i>Cocconeis neodiminuta</i> Krammer	2.2	46
<i>Cocconeis neothumensis</i> Krammer	2.65	54
<i>Cymbopleura inaequalis</i> (Ehrenberg) Krammer	0.13	5
<i>Diploneis elliptica</i> (Kützing) Cleve	0.12	9
<i>Diploneis oblongella</i> (Nägeli ex Kützing) Cleve-Euler	0.44	16
<i>Diploneis oculata</i> (Brébisson) Cleve	0.29	11
<i>Diploneis pseudovalis</i> Hustedt	0.96	37
<i>Fallacia subhamulata</i> (Grunow) D.G.Mann in Round. R.M.Crawford & D.G.Mann	0.22	17
<i>Fragilaria brevistriata</i> Grunow in van Heurck	21.06	54
<i>Fragilaria construens</i> (Ehrenberg) Grunow	3.11	37
<i>Gyrosigma acuminatum</i> (Kützing) Rabenhorst	1.27	45
<i>Hippodonta costulata</i> (Grunow) Lange-Bertalot. Metzeltin & Witkowski	1.88	41
<i>Karayevia clevei</i> (Grunow) Round & Bukhtiyarova	3.24	53
<i>Lindavia balatonis</i> (Pantocsek) T.Nakov <i>et al.</i> ,	0.96	25
<i>Navicula cryptotenella</i> Lange-Bertalot in Krammer & Lange-Bertalot	0.15	11
<i>Navicula menisculus</i> Schumann	0.14	13
<i>Navicula seminuloides</i> Hustedt	9.34	52
<i>Neidium binode</i> (Ehrenberg) Hustedt	0.24	23
<i>Neidium dubium</i> (Ehrenberg) Cleve	0.15	21
<i>Neidium iridis</i> (Ehrenberg) Cleve	0.1	5

Cont.

Table 2. (cont.)

<i>Nitzschia recta</i> Hantzsch ex Rabenhorst	0.1	10
<i>Pantocsekiella ocellata</i> (Pantocsek) Kiss & Ács	1.34	43
<i>Parlibellus crucicula</i> (W.Smith) Witkowski, Lange-Bertalot & Metzeltin	0.16	21
<i>Pinnularia balatonis</i> (Pantocsek) Mills	0.13	15
<i>Pinnularia viridis</i> (Nitzsch) Ehrenberg	0.17	12
<i>Placoneis pseudanglica</i> E.J.Cox	0.23	27
<i>Pseudostaurosira parasitica</i> (W.Smith) Morales	0.86	39
<i>Sellaphora nigri</i> (De Notaris) C.E.Wetzel & L.Ector in Wetzel <i>et al.</i> ,	1.36	30
<i>Sellaphora rotunda</i> (Hustedt) C.E.Wetzel. L.Ector. B.Van de Vijver. P.Compère & D.G.Mann	0.96	26
<i>Sellaphora utermoehlii</i> (Hustedt) C.E.Wetzel & D.G.Mann in Wetzel <i>et al.</i> ,	7.62	50
<i>Sellaphora verecundiae</i> Lange-Bertalot	0.17	8
<i>Stauroneis anceps</i> Ehrenberg	0.3	30
<i>Staurosira tabellaria</i> (W.Smith) Leuduger-Fortmorel	4.34	43
<i>Staurosirella pinnata</i> (Ehrenberg) D.M.Williams & Round	11.82	55
<i>Surirella amphioxys</i> W.Smith	0.14	19
<i>Tryblionella angustata</i> W.Smith	0.1	13

Biological remains

Diatoms

Altogether about 140 taxa were distinguished during the study (Table 2); *Amphora pediculus*, *Cocconeis neothumensis*, *Fragilaria brevistriata*, and *Karayevia clevei* were not just abundant, but ever present members of the diatom assemblage. The small fragilaroid taxa (*Staurosirella pinnata*, *Staurosira tabellaria*) and small naviculoid forms like *Navicula seminuloides* and *Sellaphora nigri*, *S. rotunda* and *S. utermoehlii* were found abundant in some layers. The genus *Amphora* (*A. pediculus* and *A. copulata*) was also often present in the samples. Biofilm constituent diatoms growing on macrophytes are also often registered, e.g. *Cocconeis neothumensis* and *C. neodiminuta*. Three planktic taxa all belonging to centric diatoms, *Aulacoseira granulata*, *Lindavia balatonis* and *Pantocsekiella ocellata* significantly contribute to the assemblages. Two diatom species, *Karayevia clevei* and *Belonastrum berolinensis* are believed to be

characteristic to Lake Balaton and were common in the majority of samples. The presence of the endemic diatom *Pinnularia balatonis* Pantocsek can also be detected along the entire core.

Diatom assemblage zones – succession of Siófok Basin

Relative frequency changes of the most common taxa throughout the sequence are shown in Figure 3. In the diatom dataset, six statistically significant diatom assemblages zones (DAZ) were distinguished. DAZs are presented below according to depth:

SIO-DAZ-1 112–104 cm – Small, but robust fragilaroid taxa (mainly *Staurosira pinnata*) with the motile *Caloneis schumanniana* and *Neidium iridis* indicate an unstable and changing environment in the lower studied layer (Stenger-Kovács *et al.*, 2018).

SIO-DAZ-2 104–82 cm – *Fragilaria brevistriata*, *Aulacoseira granulata* with small celled naviculoids (like *Navicula seminuloides* and *Sellaphora utermoehlii*) with the slightly motile

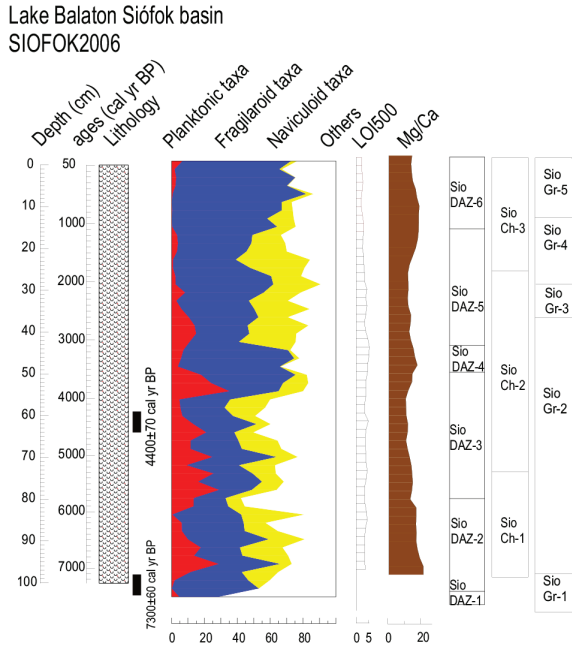


Figure 4. Lithology, Ca/Mg ratio, organic material content (LOI 550), and zone boundaries of the studied proxies in SIOFOK2006 core. Note that the diatom, geochemical and grain size analysis were carried out not on the same length of core. *Litología, relación Ca/Mg, contenido de materia orgánica (LOI 550) y límites de zona de los proxies estudiados en el núcleo SIOFOK2006. Téngase en cuenta que el análisis de diatomeas, geoquímica y tamaño de grano no se llevó a cabo en la misma longitud del testigo.*

fied *Staurosirella pinnata* is present. Increasing abundance of the epiphytic *Cocconeis neodiminuta* species is detected in this zone.

Lake level fluctuation is reconstructed based on diatom proxies reflecting the changed proportion of taxa characterized by different life forms in the assemblage. The assemblages dominated by planktic/tychoplanktic diatoms like *Aulacoseira*, *Lindavia*, *Pantocsekiella* and *Belonastrum* taxa indicate high lake levels (Fig. 4.). Fragilaroid taxa were found dominant throughout the core in the Siófok Basin due to the shallowness of Lake Balaton, but the increased ratio of planktic forms implied higher lake levels in the past. *Lindavia balatonis* was the dominant planktic species between 82–56 cm. In these layers the species *Aulacoseira granulata* was also frequently observed, exhibiting a peak between 56–52 cm.

At 46–38 cm the planktic *Pantocsekiella ocellata* reached higher abundance. According to this dataset Lake Balaton had a rather high, but fluctuating lake level before 2000 cal yr BP. Between 46–22 cm some small celled (< 10 µm) diatom taxa prevailed, suggesting a decreasing lake level. The species *Navicula seminuloides*, *Sellaphora nigri*, *S. rotunda*, and *S. utermoehlii* are also worthwhile to mention among the taxa that have been identified so far in the assemblages.

DISCUSSION

It is a well-known fact that the thickness of Holocene sediment shows a gradient along the long axis of the Lake Balaton, (e.g. Müller & Wagner, 1978; Tullner & Cserny, 2003; Visnovits, 2015): Holocene sediments can reach 10 m in the western part of the lake, where the river Zala with a rather extended catchment area enters, while it reaches only ca. 2 m in the central part of the eastern basin of the lake. Moreover, at some particular places along the lakeshore the Holocene sediments can be altogether missing (Visnovits, 2015, Visnovits *et al.*, 2015). This highly variable sediment thickness of the lake can make paleolimnological interpretation difficult.

Physical properties of SIOFOK2006 core

The relationships between LOI at 550 °C (LOI₅₅₀) and organic carbon content and between LOI at 950 °C (LOI₉₅₀) and inorganic carbon content are currently accepted as a standard proxy for organic matters (Santisteban *et al.*, 2004). Our results (very low LOI₅₅₀ and high LOI₉₅₀) are consistent with these main features, the ratio of biotas and calcareous mud of Lake Balaton.

The grain size distribution of the lake is more about diagenetic processes than the energy of water, as the sand fraction sized grains are autochthonous calcareous aggregates.

In the frame of EuLakes project four lakes were cored (two of them were shallow ones); and only Lake Balaton was found unsuitable for reconstruction of historical trophic state of the lake (<https://www.eulakes-model.eu/outputs/ecological-evolution/index.html>). Undoubtedly, the shallowness and the strong influence of wind

at Lake Balaton makes the sediment record of the lake far from an ideal subject for high-resolution paleo-ecological analyses, but our data and results obtained from SIOFOK2006 core show adequate preservation where the small, fine valves of diatoms were preserved in the mud (Kiss *et al.*, 1999).

Taxonomic remarks

Although, Lake Balaton is one of the best studied lakes in the world, the diatom flora received less attention, except in the initial pioneering phase of scientific exploration in the 19th century (Lóczy 1916, Pantocsek 1901, 1913). Unfortunately, there is no recent overview, or checklist available of the diatom flora, only a few compilations were published at the middle of the last century (Szemes 1957, Tamás 1965).

The diatom flora of Lake Balaton is unique, with some endemic species like *Pinnularia balatonis*, a ubiquitous member of the diatom assemblage throughout the core. The taxonomy of fragilaroids, exhibiting huge variability with overlapping forms has been a challenge in Lake Balaton (Ács *et al.*, 2009), especially when comparative analysis is required. It's worthwhile to mention the uncertainty of the species-level determination of *Aulacoseira* as *A. aff. granulata* in Lake Balaton. As previous paleolimnological studies of Lake Balaton revealed, the very rich diatom populations from the deeper sediment layers might help to resolve taxonomic problems. For example, the *Aulacoseira* taxa are dominant along the long axis of the lake as subfossils, a finding that helps to resolve the taxonomic difficulties concerning the genus in the lake. The species *Belonastrum berlinensis*, although a wide spread taxon (Round & Maidana 2001), is generally rare in Hungary, but here in Lake Balaton, it is regarded as a typical, frequent element of the flora.

The above mentioned taxonomical uncertainties have been known for a long time, but this study on SIOFOK2006 core offers new taxonomic challenges concerning the small naviculoid species of Lake Balaton. Among the small naviculoid (< 10 µm), species like *Navicula seminuloides*, *Sellaphora rotunda*, *S. nigrii* and *S. utermoe-hlii* have been so far successfully identified from

the assemblages. The accurate identification requiring the comparison with type materials is in progress, but still remains a challenge (as in Kiss *et al.*, 2007).

Diatom-based reconstruction on SIOFOK2006 core

Changes in the planktic/benthic ratio of diatom assemblages indicate water level fluctuation (Heinsalu *et al.*, 2007a, Korponai *et al.*, 2010). In spite of the low organic material content (LOI < 5 %), the samples prepared directly for siliceous microfossil analysis from the Holocene lacustrine sediments obtained in the Siófok Basin of Lake Balaton were suitable for diatom analysis.

Changes of water depth: Wind disturbed, very shallow water, without, or with only sporadic macrophytes can be reconstructed on the basis of the diatom assemblages from 7500 cal yr BP. The higher lake level in the Holocene was reconstructed for the period 7500-2000 cal yr BP. Gradually decreasing water depth can be inferred from 2000 cal yr BP onward.

Several studies demonstrated uneven and mosaic hydrological changes from the entire area of the historical Lake Balaton (e.g. Cserny *et al.*, (1991), Jakab *et al.*, 2005, Rokob *et al.*, 2008, Korponai *et al.*, 2010). A multiproxy, non-diatom analysis was carried out in the north-western basin of the lake reconstructing the paleoenvironment (Jakab *et al.*, 2005). This study suggested that from 5200 cal yr BP the water depth increased in the studied area. This is in line with our findings in the Siófok Basin at the south-east end of the lake. If the same phenomenon is observed in two opposite basins of the lake, then these data suggest that the lake level increase affected the entire lake and was in all probability climate-driven. In the last 2000 year history of the lake, the decreasing lake level can be due to human-induced pressure.

Mainly alkalophilous or indifferent diatom species were recorded along the sequence, indicating that there was no significant change in the water chemistry. Due to taxonomical uncertainties in Lake Balaton, and the lack of local training sets for paleoecological reconstruction, the calculation of quantitative phosphorous and pH for the

studied sequence is considered premature, but it could be an urgent task for better understanding the current ongoing processes in the lake (Korponai *et al.*, 2011). However, similar diatom records (the dominance of *Aulacoseira* taxa with small-sized fragilaroids) are often reported from large shallow lakes throughout the Holocene in Europe, maybe it is a typical assemblage for natural large lakes (Heinsalu *et al.*, 2007a).

It is well-known that the Zala river in the south-west end of the lake is the major source of inflow. The nutrient load arriving with this river controls many environmental factors in a south-west/north-east gradient along the longitudinal axis of the lake (e.g. Müller & Wagner, 1978; Tompa *et al.*, 2014), and seemingly, the preservation potential of silica in the sediment essential for the diatom record, might change from location to location also. Another high resolution study carried out in the lake is the EuLakes project (WEB1 2018), with sampling sites in the Keszthely Basin, which is exactly at the opposite end of Lake Balaton compared to the present site. Two ca. 70 cm long cores were collected at a depth of ca. 3 m. Even though the Keszthely Basin is a sub-basin of Lake Balaton receiving the highest nutrient load from the nearby Zala river during the past few decades, it is still characterised by high trophic level. However, their observation of “no diatom counts could be performed in the cleaned and mounted sub-samples due to the extremely scarce number of frustules and to their pronounced erosion. Only a few among the most robust and highly silicified taxa, could be identified in the observed samples. Planktic taxa are almost completely missing, with the exception of very rare specimens of *Cyclotella ocellata* Pantocsek (*Pantocsekiella ocellata* in this study) and *C. radiosa* (Grunow) Lemmermann / *balatonis* Pantocsek, (*Lindavia radiosa* and *L. balatonis* in this study) while the most represented benthic form is *Fragilaria martyi* (Héribaud) Lange-Bertalot” is of limited value concerning the entire lake. The conclusion of the EuLakes project that “sediment diatoms of Lake Balaton are not representative of the living community, as a result of selective diatom elimination. The highly mineral and perturbed sediments of Lake Balaton likely represent a very aggres-

sive environment for the most slender and delicate benthic and planktic taxa” can be easily refuted by the reasonably well preserved diatoms found in the Siófok Basin as presented here.

CONCLUSIONS

Several earlier studies demonstrated uneven and mosaic hydrological changes of the entire area of the historical Lake Balaton during the late glacial and Holocene times.

The use of diatoms in the paleoecological reconstruction of Lake Balaton has been questioned based on several less successful attempts conducted in the last decades. However, the SIOFOK2006 core yielded exploitable diatom flora for paleoenvironmental reconstruction.

Based on the diatom assemblages of the SIOFOK2006 core the following paleoenvironment can be reconstructed for the Siófok Basin of Lake Balaton, a wind disturbed, very shallow water without, or with only sporadic macrophytes living at the bottom of the lake. This environment prevailed from the basal layers of the core (from ca. 7500 cal yr BP) to the present, while gradually decreasing water depth can be inferred from 2000 cal yr BP onward.

Finding a high ratio of small celled naviculoid forms in the upper part of this sequence is unique, as no other earlier studied cores contained these forms in such high abundance.

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